Miocene paleoaltimetry of the Mt. Everest region

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The elevation history of the Mount Everest region is key for understanding the tectonic history of the world's highest mountain range, collisional tectonics, orogenic plateau development, onset of summer Asian monsoon, global climate change and biotic changes in Central Asia. The evolution of topography represents a complex function of regional and local climate, as well as surface and tectonic processes and it is inherently difficult to reconstruct. However, stable isotope paleoaltimetry, based on the systematic relationship between the oxygen ($\delta^{18}O$) and hydrogen (δD) isotope compositions of rainfall and elevation, permits to address this challenge of coupled tectonic and climatic research.

Most paleoaltimetry studies focus on geologic archives in sedimentary basins and to some extent neglect the role of orogen-scale deformation zones in the overall long-term landscape history. In absence of sedimentary deposits within the highly erosive Himalaya, we use the hydrogen isotope ratios (δD) of hydrous minerals that crystallized in the South Tibetan detachment (STD) shear zone during deformation (Fig. 1).

Synkinematic biotites collected systematically over 200 m of structural section from the STD into the underlying mylonitic footwall at Rongbuk Valley reveal a very constant pattern of mid Miocene 40 Ar/ 39 Ar plateau ages. The same minerals exchanged isotopically at high temperature with D-depleted water (δD_{water} = -150 ±5 ‰ equivalent to $\delta^{18}O_{water}$ = -20.0 ±1.0 ‰) that originated as high-elevation precipitation and infiltrated the crustal hydrologic system at that time [1].

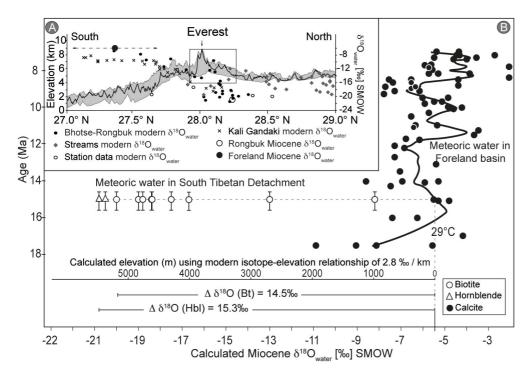


Figure 1: (A) Present-day $\delta^{18}O_{water}$ values along an elevation profile across Mount Everest. Grey area represents the elevation of a 20 km wide swath along the section with a clear $\delta^{18}O_{water}$ –elevation relationship (-2.8 %0/km) which forms the basis of our analysis. Present-day precipitation collected at Rongbuk ($\delta^{18}O_{water} \sim$ -21%0) is similar to the lowest Miocene (ca. 15 Ma) $\delta^{18}O_{water}$ values indicating similar elevations at that time.

paleoaltimetry Miocene (B) reconstruction of Mount Everest. Calculated meteoric water compositions at high elevation (South Tibetan Detachment) have distinctly lower $\delta^{18}O_{water}$ values than ageequivalent rainfall at low elevation (Foreland basin). This difference in $\delta^{18}O_{water}$ of ca. 14.5 ‰ is consistent with a Miocene paleoelevation of the Everest region in excess of 5000 m. Foreland Basin data from [3]; [4]; [5]; South Tibetan detachment: data from

Novel to our approach is to eliminate possible climate impacts on our paleoelevation estimate by comparing our high elevation hydrogen isotope record from the STD to age-equivalent oxygen isotope ratios within pedogenic carbonate from near sea level Siwalik foreland basin paleosols that record Miocene rainfall conditions in the

Himalayan foothills (Fig. 1). The relative difference between mid Miocene meteoric water compositions in the Siwalik foreland ($\delta^{18}O_{water} \approx -5.5 \pm 1.0\%$) based on $\delta^{18}O$ of pedogenic calcite) and the high elevation STD ($\delta^{18}O_{water} \approx -20 \pm 1\%$) based on δD of biotite) is $\Delta \delta^{18}O_{water} \approx 14.5 \pm 1.0\%$ which is consistent with mean elevations of ~ 5200 m for the Everest region at ~ 15 Ma.

These similar-to-modern Miocene mean elevations have various implications: 1) If Himalayan topography indeed plays a pivotal role in controlling monsoonal rainfall patterns on the Indian subcontinent [2], presence of a high Himalayan range at ~ 15 Ma indicates strengthening of the Asian summer monsoon earlier than previously thought; 2) Mid Miocene high elevation Himalayan topography must have influenced mid Miocene rainfall patterns on the Tibetan plateau including the presence of a strong Himalayan rain shadow.

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