

PHOTOPLATES 5 - 8

PLATE 5

1. The Lobe B deposits of the proximal lobe zone are composed of thick-bedded, laterally extensive sandstones (person for scale). The deposits are highly amalgamated and characterised by a high net sand content. Note the pervasive covering by lichens.
2. Water-escape features, such as these burst-through features (sand volcanoes) which heavily deform the original traction carpet lamination, are common in Lobe B deposits of the proximal lobe zone (lens cap for scale at bottom left).
3. A coarse-grained sandstone ($S_{1,2}$ of Lowe 1982) with traction structures is erosionally cut into by a granule sandstone (hammer for scale). The presence of an escape burrow (arrow) burrowing through both beds suggests at times rapid sedimentation rates within lobe B depositional environment.
4. Massive composite sandstone are common within the lobe B deposits (hammer for scale). With their often poorly distinguished amalgamation surfaces, the lack of grading and an abundance of floating shale clasts they resemble deep-water massive sandstones (DWMS) of Stow & Johansson (2000).
5. Within the highly amalgamated lobe B deposits wedging beds (arrow) feature rarely (hammer for scale). They are believed to compensate topographic expression of previously deposited deposits. More often small-scale wedging geometries are associated with the depositional wings of distributary channels.
6. Detail of the offset-stacked distributary channels near the top of the proximal lobe zone (hammer for scale). Distributary channel (2) cuts down into the stacked channel-fill deposit of distributary channel (1). The channels were too shallow to contain volumetrically larger flows which resulted in “wing-like” deposition outwith the channel margins (3). Selective erosion may account for the stepped channel margin (arrow).
7. Thin-bedded, medium to fine sand-sized sandstones ($T_{(b)c-d,e}$ of Bouma 1962) and interbedded shales represent the finest fraction of the lobe fringe deposits of the proximal lobe zone (pen for scale). Their occurrence is rare, occasionally separating lobe B packages. They are characterised by a low sand:shale ratio (2:1).

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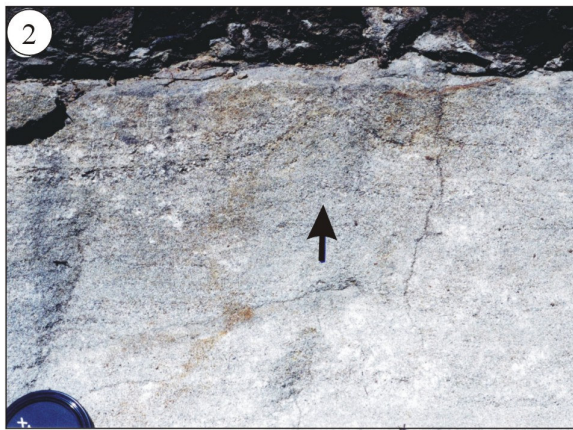


PLATE 6

1. The Lobe C deposits of the distal lobe zone are characterised by laterally extensive, parallel bedded sandstones (scale: 1 m). The sandstones are typically composed of coarse- to medium-grained sandstones, separated by a few cm of thick fine sandstone and shale intercalations. The net sand content is in average greater than 80 %.
2. Sharp-based sandstone with massive bases, uniformity in grain size nearly throughout the whole bed and only rapid fining at the top are commonly observed (pen for scale). They can be classified as S_{2-3} (Lowe 1982) or T_{a-d} (Bouma 1962). Note the faintly aligned, weathered shale chips at the top right and burrow at the bottom right (arrow). Both features are commonly found within the Lobe C sediments.
3. Excellent preservation of bioturbation on bedding surfaces can be observed in the lobe C deposits. *Zoophycus* is a commonly found trace fossils (lenscap for scale).
4. Slumping, especially involving sandstone beds, is rarely observed within the distal lobe environment (scale: 1 m). The shown slump package (arrow; slump axis pointing to left) is composed of medium sand-sized beds intercalated with shales and fine sand typical of lobe fringe deposits. As semi-lithified sediments they may have collapsed into an intervening depression (interlobe environment?) The sharp onset of thick, relatively coarse lobe deposits suggests a sudden shift of lobe sedimentation into this particular area.
5. The groove casts at the sole of this bed show a 30° divergence in transport direction, pointing to a general NE (25 - 55°) direction (hammer for scale).
6. This lobe fringe package is marked by a distinct coarsening- and thickening-upward sequence believed to result from lobe migration into this particular area. The net sand content increases from approximately 50 to 75 %.
7. The fine-grained interlobe deposits are characterised by a low net sand content (~ 60 %). They are composed of thin-bedded, fine-grained sandstones ($T_{c-d,e}$ of Bouma 1962; hammer for scale). The presence of the thick, isolated sandstone bed suggests that occasionally larger turbiditic flows reached into this otherwise relatively quiet depositional environment.
8. The top of the Seyhan River section is characterised by thick, monotone successions of fan fringe deposits (person for scale). Turbiditic flows capable of depositing thicker and coarser sandstone beds (arrow) only occasionally reach this area.

PLATE 6

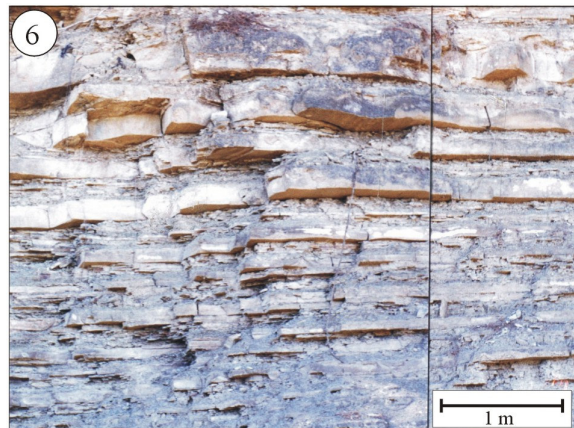


PLATE 7

1. Top of matrix-supported G5 deposit with large marl (1) and smaller limestone clasts. Small, deformed shale clasts are roughly aligned, some syndimentary deformation (loading? arrow) is present (14/19-21: 10573 ft).
2. Sand-shale, matrix-rich debris flow deposits (G6, bottom) overlain by clean gravelly, limestone clast-rich conglomerate (G3, top). G3 loading into G6 (arrow). Maximum clast size G6 is 10 cm and G3 is 2.5 cm. G6 facies contains shell fragments (arrow; 14/19-E7: 9856 – 9855 ft).
3. Well cemented GS2 deposit with distinct pebble-delineated traction carpets (1). Reverse grading, i.e. upward increase in pebble size. Winnowing of sand along traction surface (?). Succeeding flow eroded into GS facies (arrow; 14/19-21: 10592.5 ft).
4. GS3 facies with abundant aligned, subangular shale rip-up clasts and smaller limestone pebbles in coarse sand matrix. Some shale clasts are thoroughly bioturbated (1). Overlying sandstone erodes into GS3 deposit (arrow; 14/19-E7: 9891 ft).
5. Poorly sorted, pebble-rich area shows preferential carbonate cementation (nodule; 14/19-20: 9869 ft).

PLATE 7

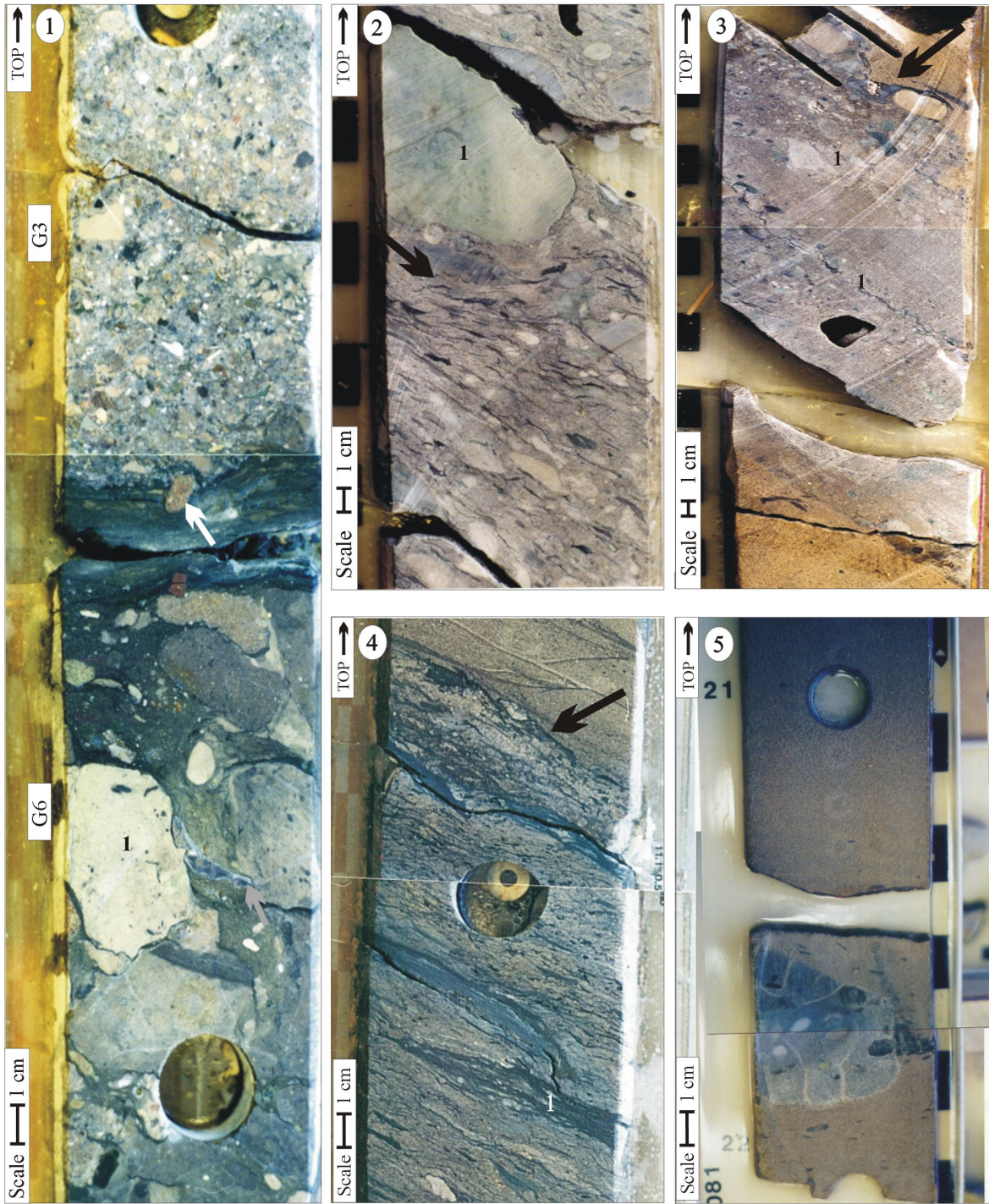


PLATE 8

1. Disorganised GS5 deposit with reworked carbonate nodules (1) and shell fragments (2: gastropod; 14-19-E3: 10668 ft)
2. Medium-grained, massive S1 sandstone (14/19-E2: 9134.5 ft).
3. Medium-grained, well laminated S3 sandstone cut into (arrow) underlying fine-grained, well laminated top of S2 sandstone (14/19-E2: 9099 ft).
4. Bimodal SM1 facies with reddish shale matrix and granule to small pebble-sized grains. The shale content is greater than 15% (14/19-E2: 9075 ft).
5. Succession of thin-bedded SH2-S3-SH2-SM3-SM2 deposits(14/19-E4: 9685 ft). Erosive contacts are common (arrows), nodule formation (N) rare.
6. Highly bioturbated marls (hemipelagic Bi facies) interbedded with SH1 facies (top) and highly bioturbated thin-bedded turbidites of SM3 lithofacies type.(14/19- E2: 9163 ft).

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